

On Wind-Driven Electrojets at Magnetic Cusps in the Nightside Ionosphere of Mars

M. O. Fillingim¹, R. J. Lillis¹, S. L. England¹,
L. M. Peticolas¹, D. A. Brain¹, J. S. Halekas¹, C. Paty²,
D. Lummerzheim³, and S. W. Bougher⁴

¹Space Sciences Laboratory, University of California, Berkeley

²School of Earth and Atmospheric Sciences, Georgia Institute of
Technology, Atlanta

³Geophysical Institute, University of Alaska, Fairbanks

⁴Department of Atmospheric, Oceanic and Space Sciences,
University of Michigan, Ann Arbor

Outline

Summary

Martian Magnetic Fields and Cusps

Martian Ionospheric Dynamo

Martian Ionospheric Currents

Terrestrial Auroral Electrojets

Martian Auroral Electrojets

Variability

Caveats/Assumptions/Simplifications

Summary

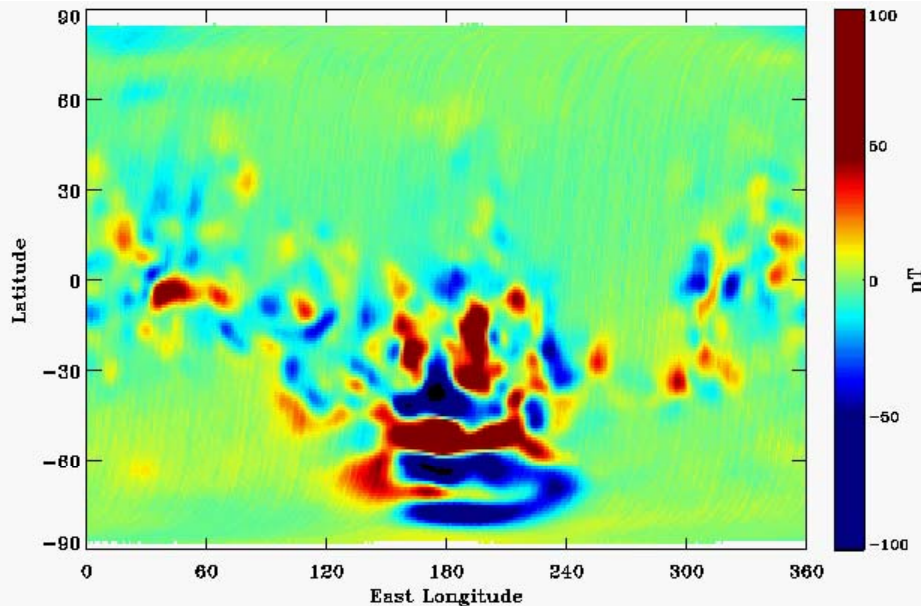
Summary

- The complex magnetic topology at Mars allows solar wind (and accelerated) electrons to ionize the nightside atmosphere in limited regions (cusps) forming a patchy nightside ionosphere
- Neutral winds drive ionospheric currents at altitudes where ions are collisionally coupled to the neutral atmosphere while electrons are magnetized → dynamo region
- Inhomogeneities in the ionospheric conductivity lead to polarization electric fields and secondary ionospheric currents – secondary currents can reinforce original currents forming electrojets
- The magnetic signatures of electrojets can be measured from orbit and from the surface

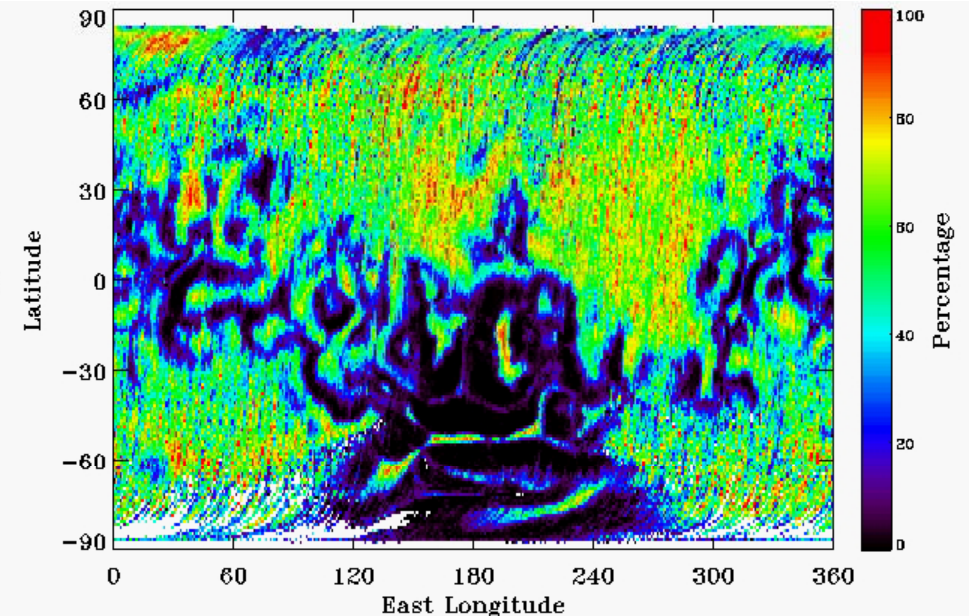
Martian Magnetic Fields and Cusps

- No global magnetic field ***but*** strong crustal fields
- Cusps form where radial crustal fields connect to the IMF
→ solar wind has access to the atmosphere → ionization
- Non-uniform global distribution of cusps and ionization
- Accelerated electrons, ionospheric structure, and aurora associated with cusps

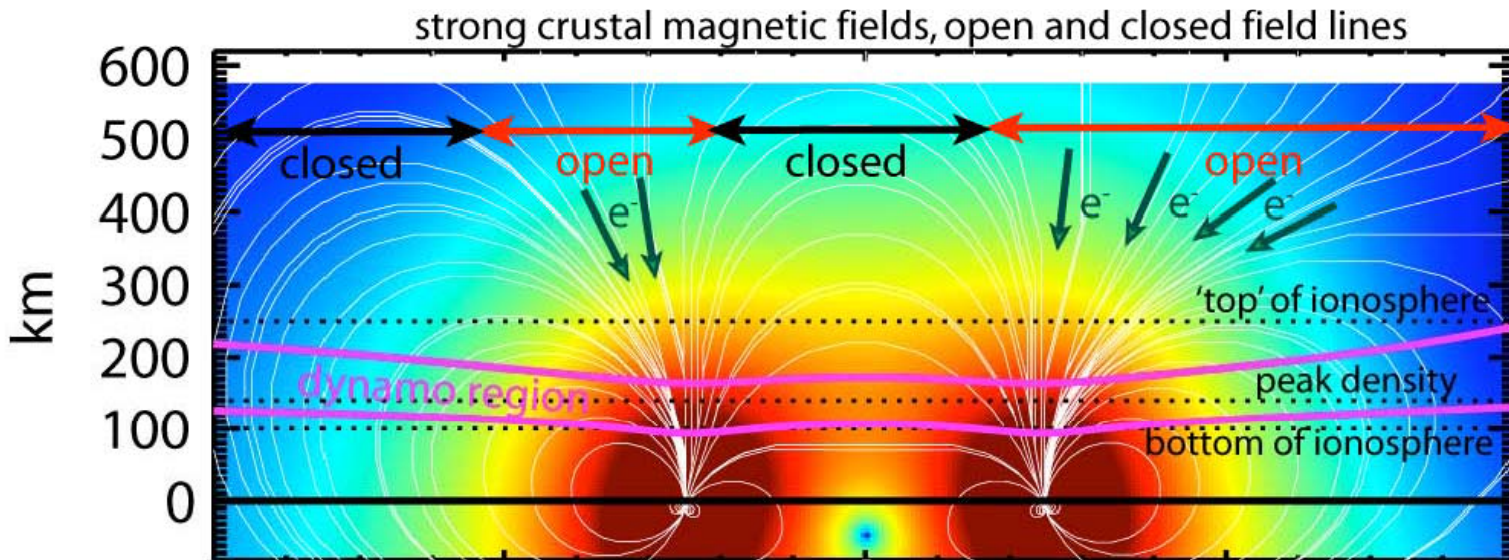
Radial component of B



Probability of observing loss cones on the nightside



Martian Ionospheric Dynamo

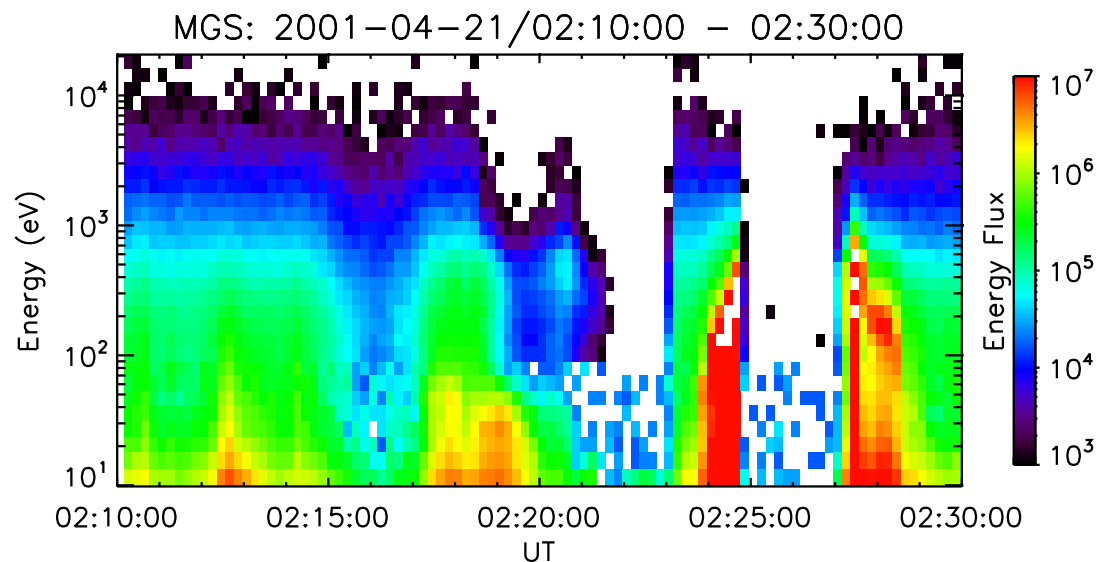


- Ionospheric currents exist where ions are collisional ($\Omega_i < \nu_{in}$) but electrons are magnetized ($\Omega_e > \nu_{en}$) \rightarrow **dynamo region**
- Crustal magnetic fields alter the ionospheric electrodynamics
- The altitude of the dynamo is geographically dependent
- Currents vary on the same spatial scales as the crustal fields do at ionospheric altitudes ($\sim 100 - 600$ km)

Martian Ionospheric Currents

(see Fillingim *et al.* (2010), *Icarus*, 206(1), pp. 112-119.)

1. Start with observed electron energy spectra



Electron spectrogram
observed by Mars
Global Surveyor on the
nightside at 400 km
Regions of accelerated
electrons (at cusps) &
“voids” of few electrons

Martian Ionospheric Currents

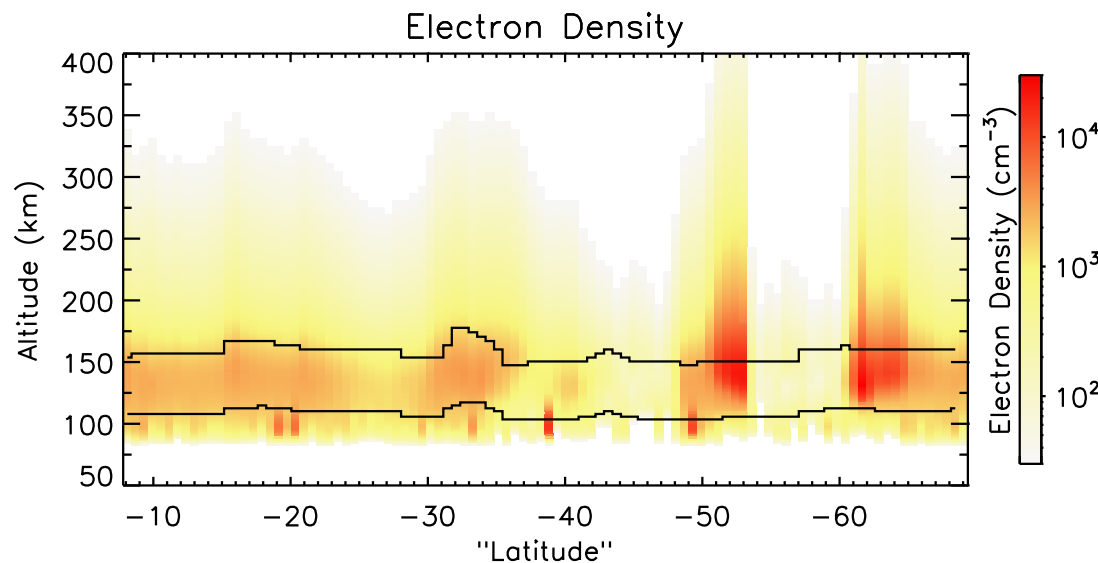
(see Fillingim *et al.* (2010), *Icarus*, 206(1), pp. 112-119.)

1. Start with observed electron energy spectra
2. Calculate ionization rate
 - neutral atmosphere (MTGCM) of Bougher *et al.* [2009]
 - electron transport code of Lummerzheim & Lilensten [1994]
 - magnetic field model of Cain *et al.* [2003] (path length)

Martian Ionospheric Currents

(see Fillingim *et al.* (2010), *Icarus*, 206(1), pp. 112-119.)

1. Start with observed electron energy spectra
2. Calculate ionization rate
3. Compute resulting electron density, n_e
 - assume photochemical equilibrium, i.e., $n_e(z) = \sqrt{P(z)/\alpha_{\text{eff}}(z)}$
 - assume all ions are O_2^+ , $\alpha_{\text{eff}}(z)$ is O_2^+ recombination rate
 - electron temperature, T_e , is equal to measured daytime T_e



Computed n_e versus altitude and latitude

Black lines bound dynamo region → currents coincide with ionospheric peak

Martian Ionospheric Currents

(see Fillingim *et al.* (2010), *Icarus*, 206(1), pp. 112-119.)

1. Start with observed electron energy spectra
2. Calculate ionization rate
3. Compute resulting electron density, n_e
4. Add external force \rightarrow neutral winds ($u_x = 100$ m/s northward)

Martian Ionospheric Currents

(see Fillingim *et al.* (2010), *Icarus*, 206(1), pp. 112-119.)

1. Start with observed electron energy spectra
2. Calculate ionization rate
3. Compute resulting electron density, n_e
4. Add external force \rightarrow neutral winds ($u_x = 100$ m/s northward)
5. From equations of motion, calculate particle velocities, $\mathbf{v}_{i,e}$

$$-1/n_{i,e} \nabla(n_{i,e} kT_{i,e}) + m_{i,e} \mathbf{g} + q(\mathbf{E} + \mathbf{v}_{i,e} \times \mathbf{B}) - m_{i,e} \nu_{in,en} (\mathbf{v}_{i,e} - \mathbf{u}) = 0$$

pressure gravity electric magnetic collisions with
gradient field field neutrals

Martian Ionospheric Currents

(see Fillingim *et al.* (2010), *Icarus*, 206(1), pp. 112-119.)

1. Start with observed electron energy spectra
2. Calculate ionization rate
3. Compute resulting electron density, n_e
4. Add external force \rightarrow neutral winds ($u_x = 100$ m/s northward)
5. From equations of motion, calculate particle velocities, $\mathbf{v}_{i,e}$

$$-1/n_{i,e} \nabla(n_{i,e} kT_{i,e}) + m_{i,e} \mathbf{g} + q(\mathbf{E} + \mathbf{v}_{i,e} \times \mathbf{B}) - m_{i,e} \nu_{in,en} (\mathbf{v}_{i,e} - \mathbf{u}) = 0$$

pressure gradient gravity electric field magnetic field collisions with neutrals

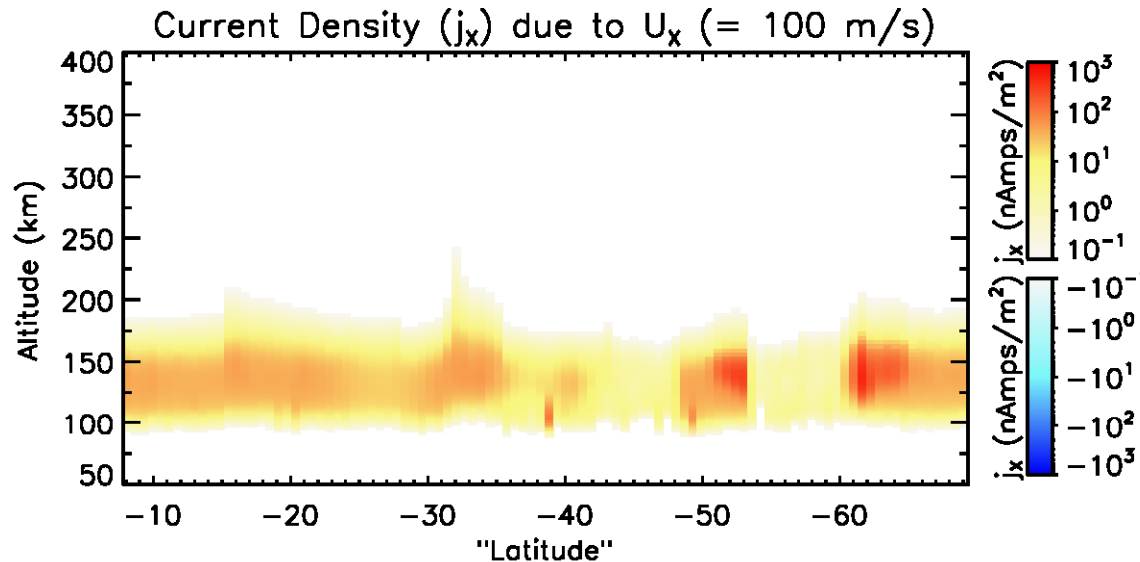
* Assume $\mathbf{B} = B_z$ and $\mathbf{u} = u_x$

Martian Ionospheric Currents

(see Fillingim *et al.* (2010), *Icarus*, 206(1), pp. 112-119.)

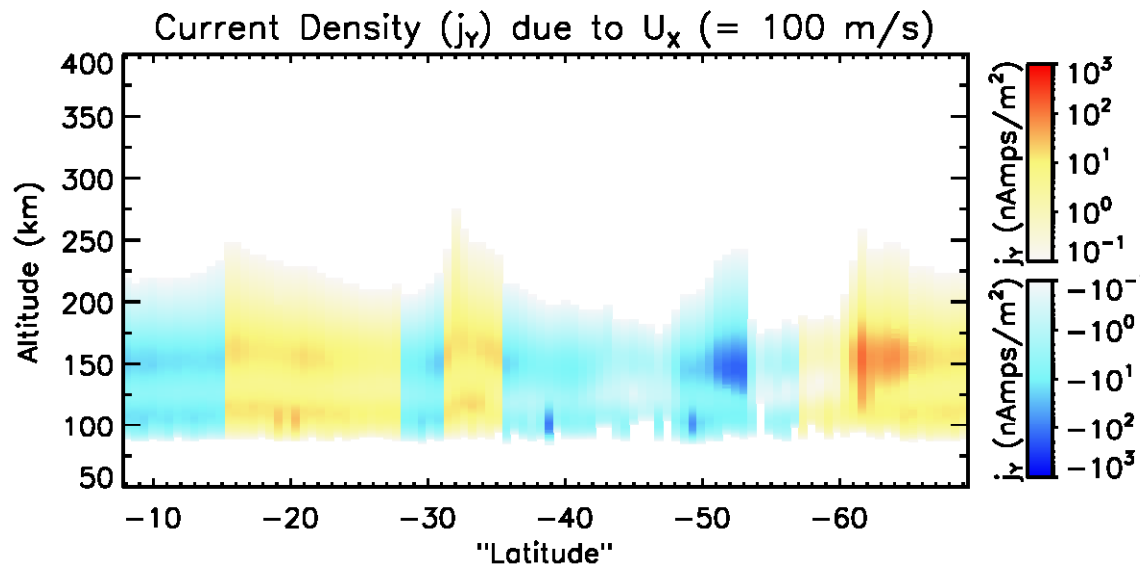
1. Start with observed electron energy spectra
2. Calculate ionization rate
3. Compute resulting electron density, n_e
4. Add external force \rightarrow neutral winds ($u_x = 100$ m/s northward)
5. From equations of motion, calculate particle velocities, $\mathbf{v}_{i,e}$
6. Calculate currents: $\mathbf{j} = nq(\mathbf{v}_i - \mathbf{v}_e)$

Martian Ionospheric Currents



X- (north) and Y- (west) components of currents driven by a uniform northward neutral wind; $u_x = +100 \text{ m/s}$

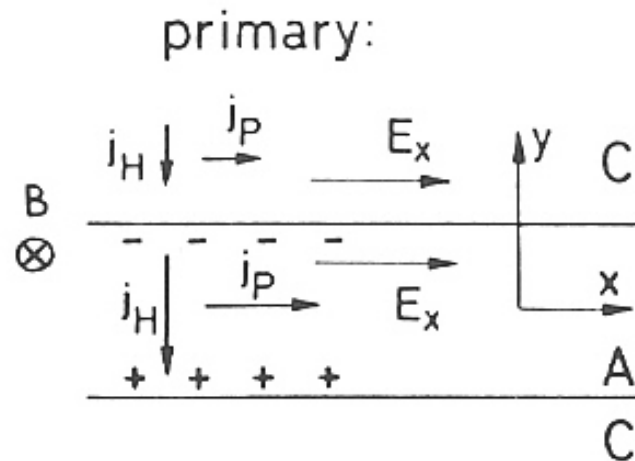
Collisional ions carry northward currents, j_x



East-west currents, j_y , carried by magnetized electrons as they drift in the $-\mathbf{F} \times \mathbf{B}$ direction where $\mathbf{F} = F_x = m_e v_{en} u_x$

What about **secondary effects...**?

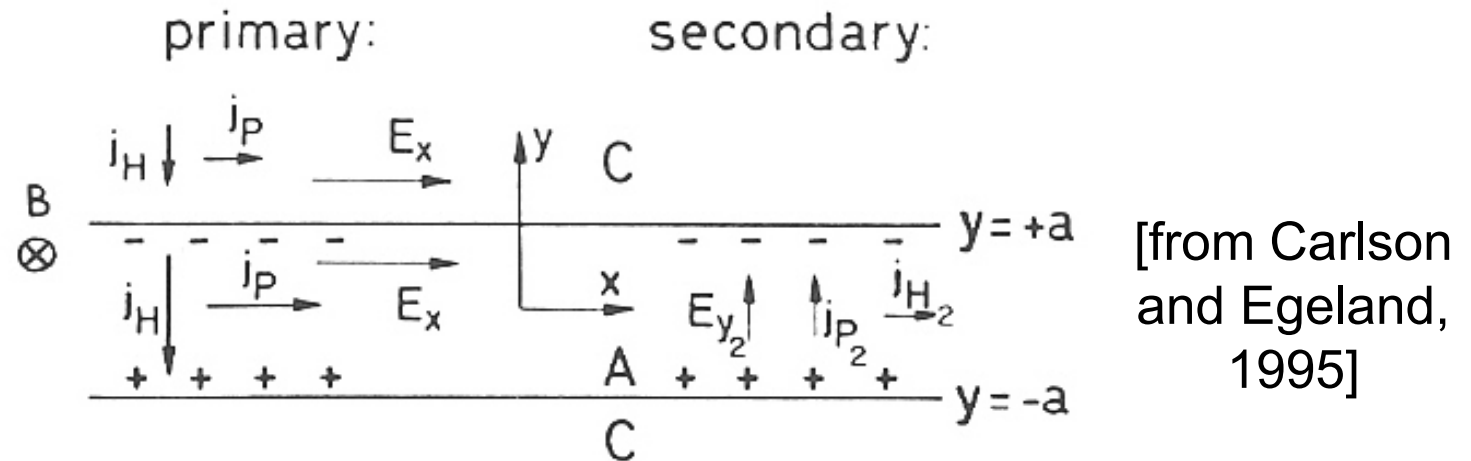
Terrestrial Auroral Electrojets



[from Carlson
and Egeland,
1995]

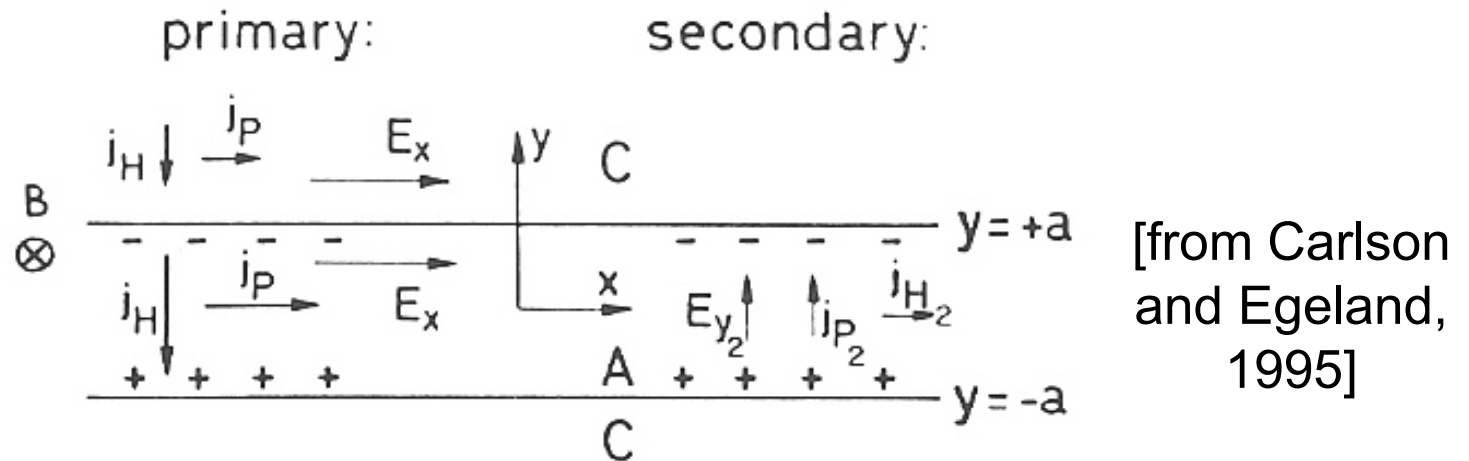
- Particle precipitation (i.e., aurora) can create a channel of enhanced ionization and enhanced conductivity, σ (region A)
- An external force (electric field, \mathbf{E} ,) drives ionospheric currents
- Collisional ions carry current parallel to \mathbf{E} : Pedersen current, j_P
- Magnetized electrons carry current perpendicular to both \mathbf{E} and \mathbf{B} : Hall current, j_H
- Currents in region A are stronger due to higher conductivity \rightarrow difference in j_H leads to charge accumulation at the edges of A

Terrestrial Auroral Electrojets



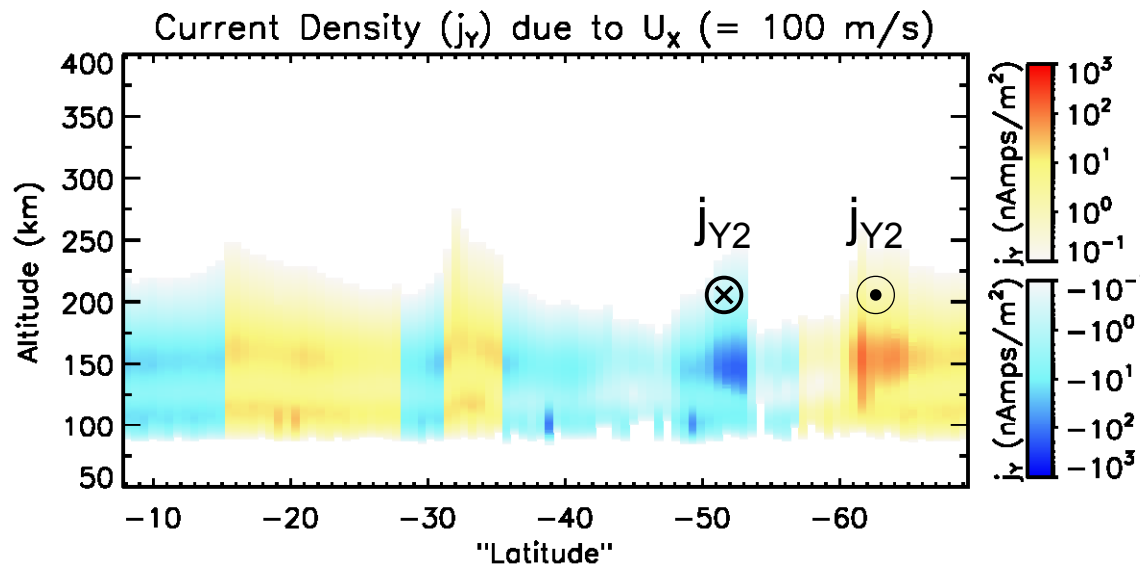
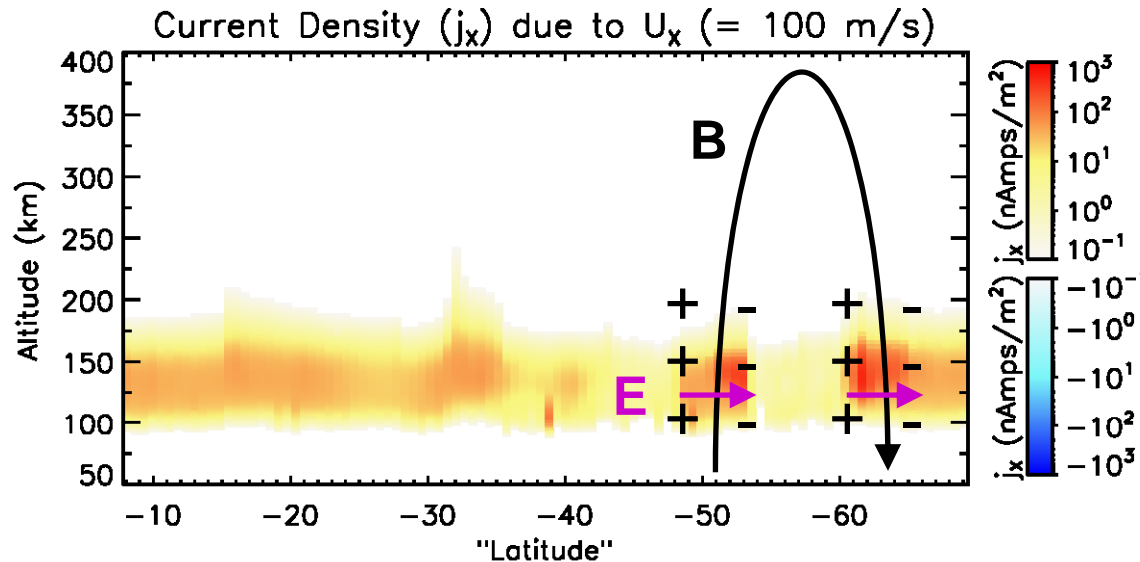
- Charge separation creates a secondary electric field, E_{y2} , which drives secondary Pedersen and Hall currents, j_{P2} and j_{H2}
- j_{P2} partially cancels j_H in region A
 → current continuity in y-direction across A-C boundary
- j_{H2} adds to j_P enhancing original current → **electrojet**
- Can an analogous situation occur in the nightside ionosphere of Mars...?

Terrestrial Auroral Electrojets



- Charge separation creates a secondary electric field, E_{y2} , which drives secondary Pedersen and Hall currents, j_{P2} and j_{H2}
- j_{P2} partially cancels j_H in region A
 → current continuity in y-direction across A-C boundary
- j_{H2} adds to j_P enhancing original current → **electrojet**
- Can an analogous situation occur in the nightside ionosphere of Mars...?
- **Yes – in magnetic cusps!**

Martian Auroral Electrojets



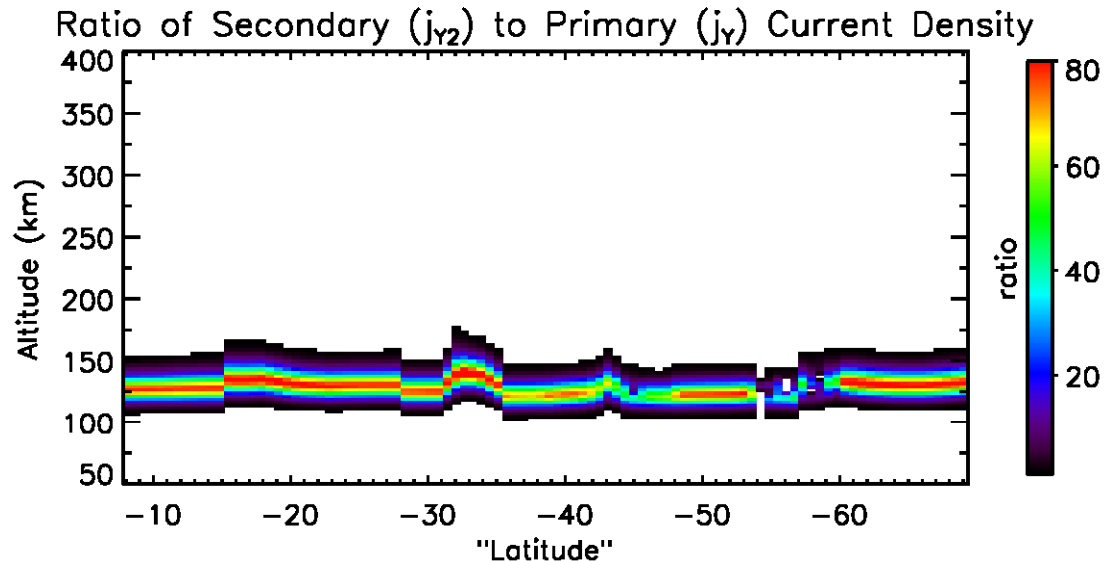
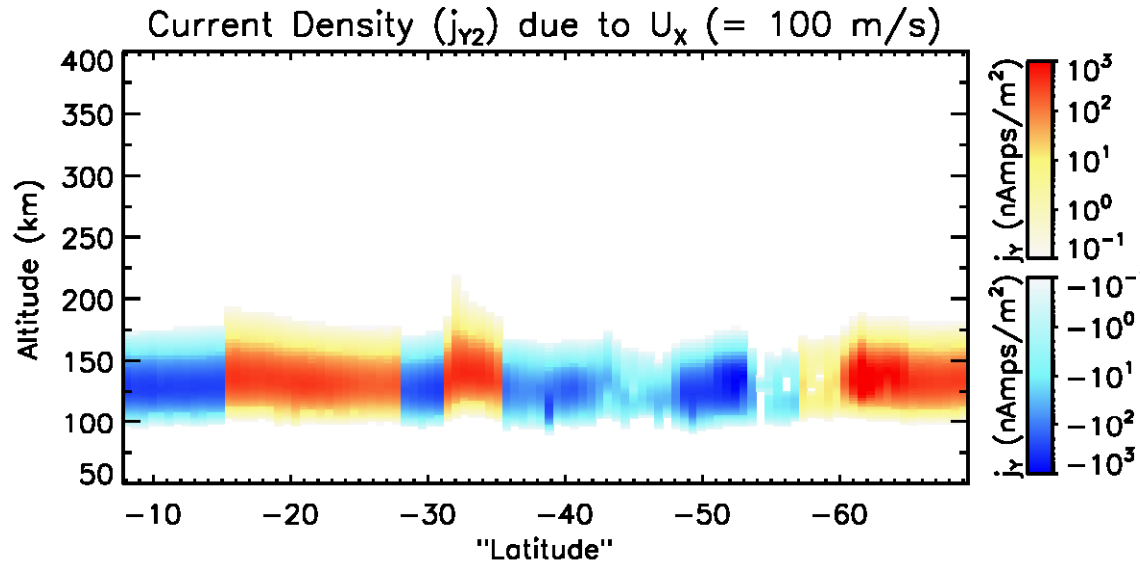
Variations in spectra of precipitating electrons
 → variations in n_e
 → variations in σ
 → variations in j

Neglecting parallel currents, j_x must be continuous (and small)

Charge accumulates at edges of high σ cusps
 → creates **southward E**

E drives secondary Hall currents, j_{Y2} , enhancing original j_y

Martian Auroral Electrojets



Secondary east-west
Hall current, j_{Y2} ,
calculated assuming j_X
continuous and $= j_X^{\min}$

$$E_x = (j_X - j_X^{\min}) / \sigma_P,$$

$$j_{Y2} = E_x \sigma_H \approx j_X \sigma_H / \sigma_P$$

Current increase, j_{Y2}/j_Y

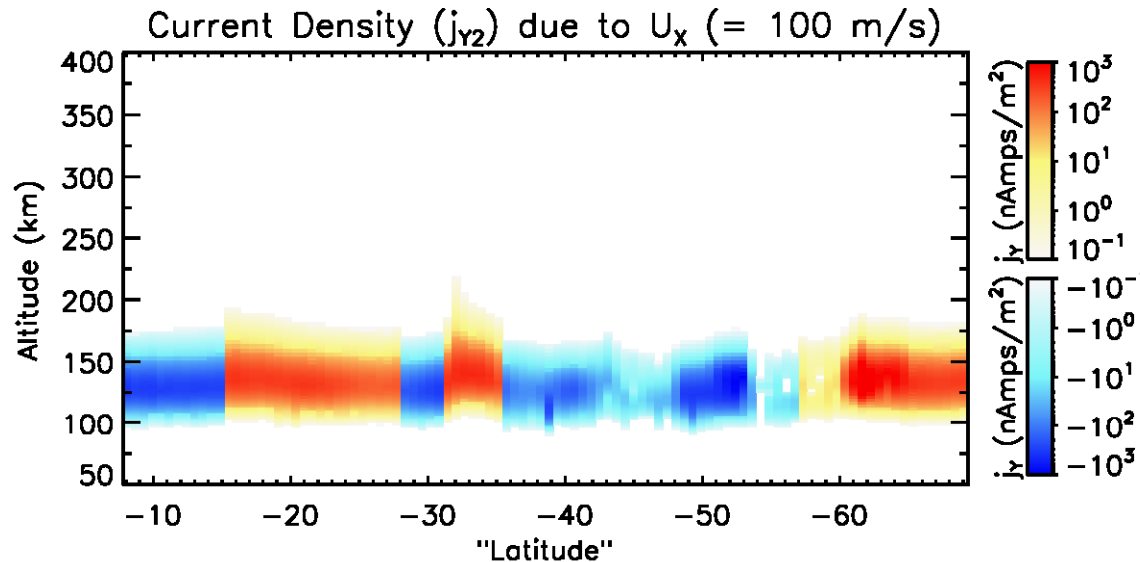
$$j_{Y2}/j_Y \approx j_X/j_Y \sigma_H / \sigma_P$$

$$\approx [\sigma_H / \sigma_P]^2$$

Factor of ~ 80 increase
 \rightarrow **electrojets**

Large j_X , large increase

Martian Auroral Electrojets



Secondary east-west Hall current, j_{Y2} , calculated assuming j_X continuous and $= j_X^{\min}$

$$E_x = (j_X - j_X^{\min}) / \sigma_P,$$

$$j_{Y2} = E_x \sigma_H \approx j_X \sigma_H / \sigma_P$$

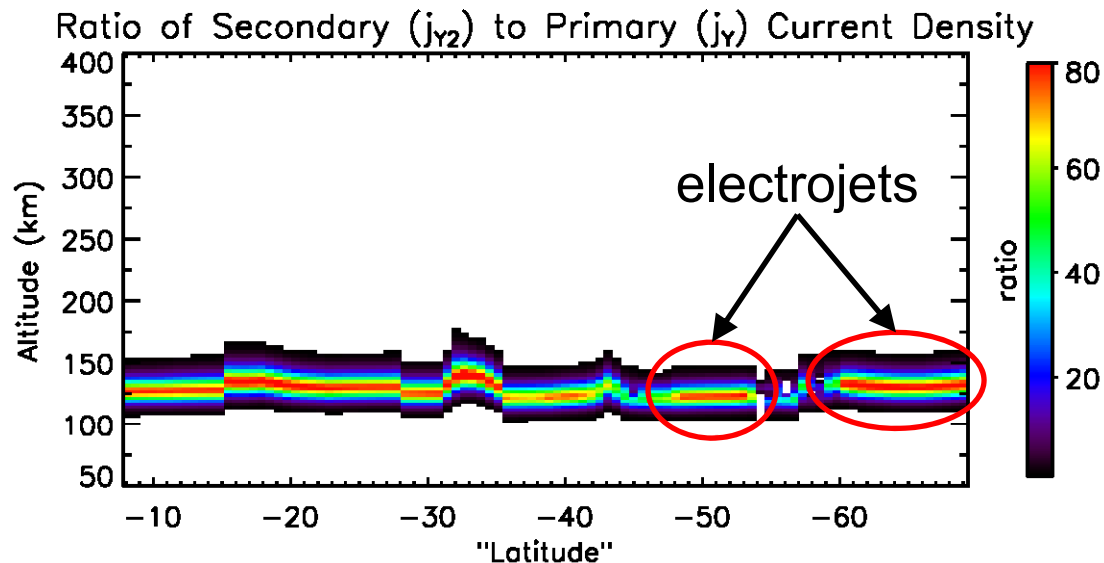
Current increase, j_{Y2}/j_Y

$$j_{Y2}/j_Y \approx j_X/j_Y \sigma_H / \sigma_P$$

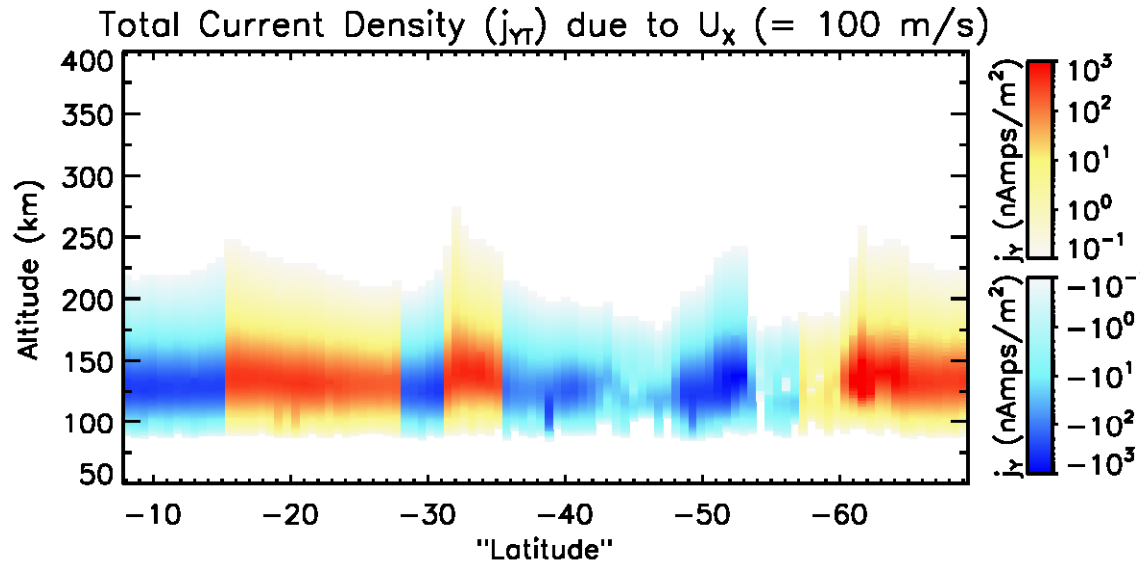
$$\approx [\sigma_H / \sigma_P]^2$$

Factor of ~ 80 increase
→ electrojets

Large j_X , large increase



Martian Auroral Electrojets

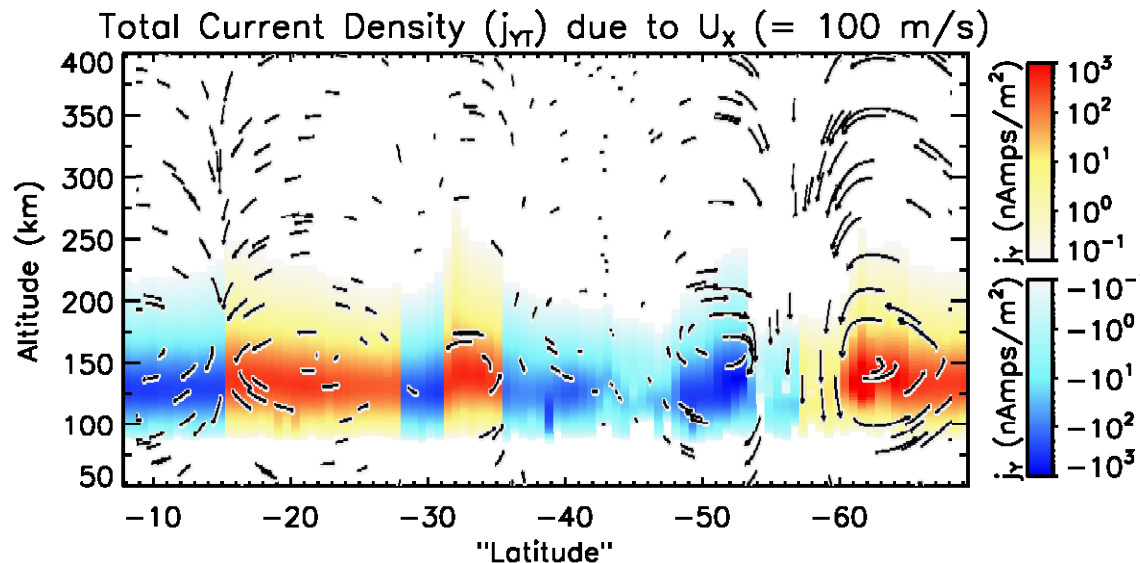


Total current density

$$j_{YT} = j_Y + j_{Y2}$$

Solve Biot-Savart Law
to find ΔB due to j_{YT}

Max j_{YT} at -50° & -65° ;
max ΔB in region
between $-50 - -65^\circ$



at 400 km, $\Delta B \approx 10 \text{ nT}$,
 $B_{\text{ambient}} \approx 100 \text{ nT}$ (10%)

at 150 km, $\Delta B \approx 50 \text{ nT}$,
 $B_{\text{ambient}} \approx 500 \text{ nT}$ (10%)

at surface, $\Delta B \approx 10 \text{ nT}$,
 $B_{\text{ambient}} > 1000 \text{ nT}$ (1%)

Variability

- Wind driven electrojets are variable; periodic changes in **conductivity gradients** and neutral wind **speed** and **direction** affect intensity of electrojets
- Diurnal:
 - In **sunlight**, conductivity gradients are weaker (solar EUV) → j_x is “more continuous” → weaker electrojets
 - Wind patterns change with **local time** [Bougher *et al.*, 2000] northward winds in southern hemisphere pre-midnight; westward wind post-midnight → weaker electrojets
- Seasonal:
 - Nightside wind patterns also change with **season** northward winds at equinox and southern summer solstice; eastward winds at northern summer solstice → weaker EJ

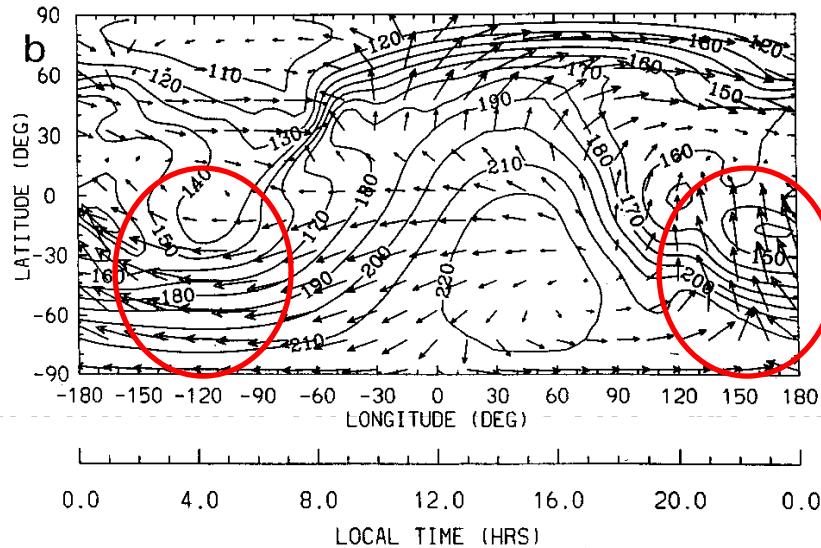
Variability

- Wind driven elec
- **conductivity gr**
- affect intensity o

changes in
ed and direction

- **Diurnal:**

- In sunlight, cc
- j_x is “more con



er (solar EUV) →

- Wind patterns change with **local time** [Bougher *et al.*, 2000]
northward winds in southern hemisphere pre-midnight;
westward wind post-midnight → weaker electrojets

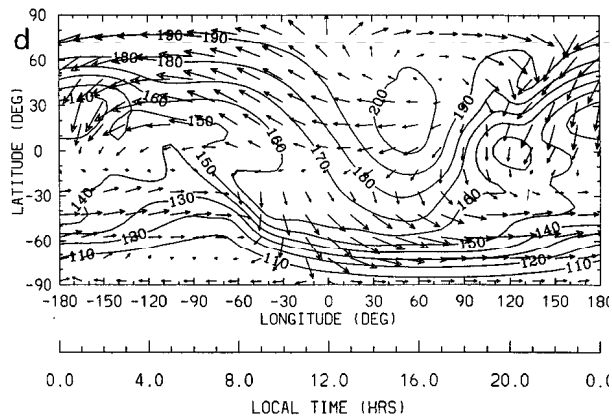
- **Seasonal:**

- Nightside wind patterns also change with **season**
northward winds at equinox and southern summer solstice;
eastward winds at northern summer solstice → weaker EJ

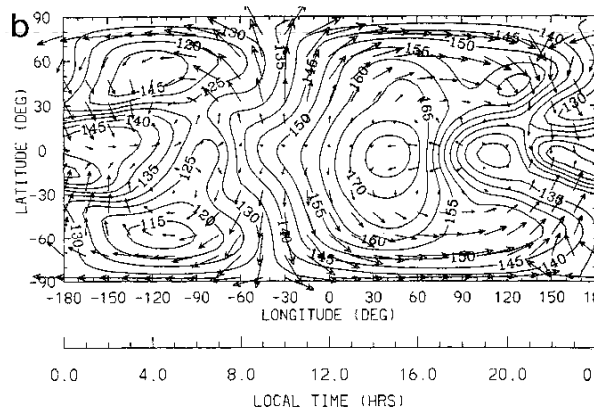
Variability

- Wind driven electrojets are variable; periodic changes in **conductivity gradients** and neutral wind **speed and direction** affect intensity of electrojets

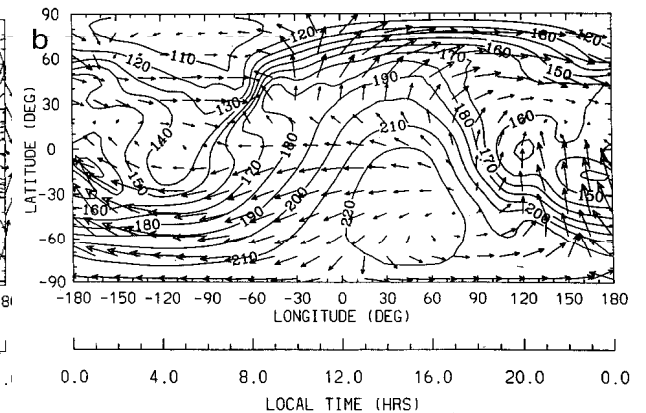
Northern Summer Solstice



Equinox



Southern Summer Solstice



- **Seasonal:**

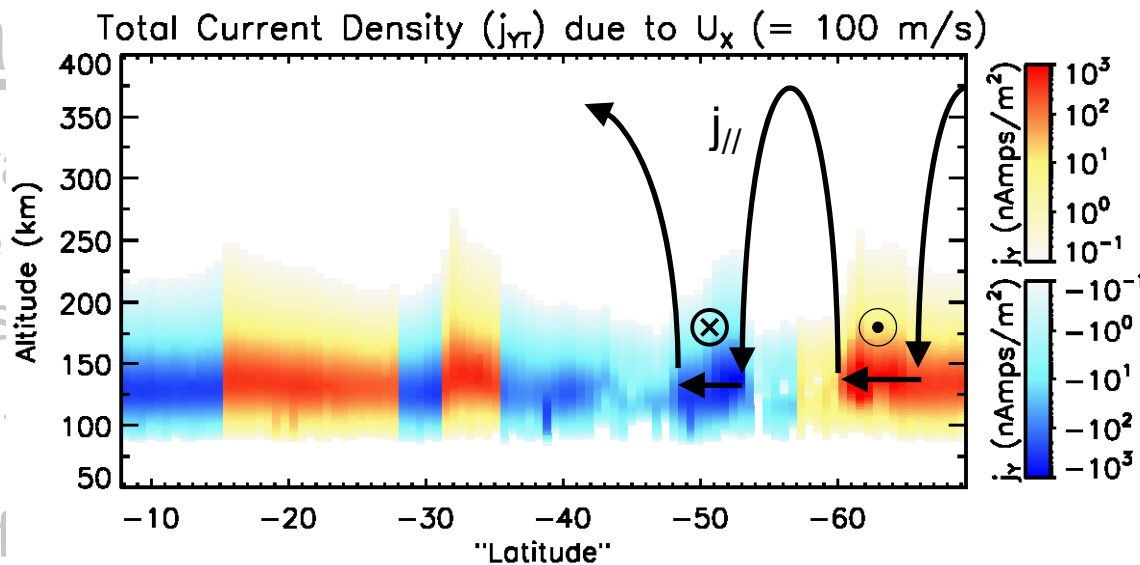
- Nightside wind patterns also change with **season**
northward winds at equinox and southern summer solstice;
eastward winds at northern summer solstice → weaker EJ

Caveats/Assumptions/Simplifications

- Electron transport code does not include magnetic gradients: straight field lines with constant magnitude and dip angle
→ Bad assumption for anisotropic electrons (Lillis *et al.*, 2009)
- For current calculations, use unrealistic geometry → $\mathbf{B} = B_z$
- Neglect effects of external (magnetospheric) electric fields
- Ignore (observed) parallel currents: $j_{\parallel} \sim 0.5 - 1 \mu\text{A}/\text{m}^2$
[Brain *et al.*, 2006; Halekas *et al.*, 2006]
 j_{\parallel} will decrease – but not nullify – magnitude of electrojets
→ 3-D current system analogous to Earth's auroral region
- Currents modify magnetic field (which modify currents...)
- What is needed to more adequately address these problems?
→ Geometrically accurate, self-consistent, 3-D model of the electrodynamics of the Martian ionosphere (see Poster 41)

Cavea

- Electron trajectory straight field → Bad assumption
- For current
- Neglect eff



assumptions

- Ignore (observed) parallel currents: $j_{\parallel} \sim 0.5 - 1 \mu\text{A/m}^2$ [Brain *et al.*, 2006; Halekas *et al.*, 2006]
 j_{\parallel} will decrease – but not nullify – magnitude of electrojets
→ 3-D current system analogous to Earth's auroral region
- Currents modify magnetic field (which modify currents...)
- What is needed to more adequately address these problems?
→ Geometrically accurate, self-consistent, 3-D model of the electrodynamics of the Martian ionosphere (see Poster 41)

gradients:
angle
(*et al.*, 2009)
→ $\mathbf{B} = B_z$
c fields

Caveats/Assumptions/Simplifications

- Electron transport code does not include magnetic gradients: straight field lines with constant magnitude and dip angle
→ Bad assumption for anisotropic electrons (Lillis *et al.*, 2009)
- For current calculations, use unrealistic geometry → $\mathbf{B} = B_z$
- Neglect effects of external (magnetospheric) electric fields
- Ignore (observed) parallel currents: $j_{\parallel} \sim 0.5 - 1 \mu\text{A}/\text{m}^2$
[Brain *et al.*, 2006; Halekas *et al.*, 2006]
 j_{\parallel} will decrease – but not nullify – magnitude of electrojets
→ 3-D current system analogous to Earth's auroral region
- Currents modify magnetic field (which modify currents...)
- What is needed to more adequately address these problems?
→ Geometrically accurate, self-consistent, 3-D model of the electrodynamics of the Martian ionosphere (see Poster 41)

Summary

- The complex magnetic topology at Mars allows solar wind (and accelerated) electrons to ionize the nightside atmosphere in limited regions (cusps) forming a patchy nightside ionosphere
- Neutral winds drive ionospheric currents at altitudes where ions are collisionally coupled to the neutral atmosphere while electrons are magnetized → dynamo region
- Inhomogeneities in the ionospheric conductivity lead to polarization electric fields and secondary ionospheric currents – secondary currents can reinforce original currents forming electrojets
- The magnetic signatures of electrojets can be measured from orbit and from the surface